VIRS Images





Fig. 1 Color composite RGB Image (Tibetan Plateau) *R: ch1, G: ch2, B: ch4*



Fig.3 Color composite RGB Image (South of the Brazil) *R: smaller BTD between ch4 and ch5 G: larger BTD between ch4 and ch5 B: brightness temperature at ch4*



Fig.2 Gray-scale image for the ratio of ch1 and ch2 (Tibetan Plateau)



Fig.4 Gray-scale image for BTD between ch4 and ch5 (South of the Brazil)

可視赤外観測装置(VIRS)によるチベット及び南ブラジルの観測

図1は1997年12月21日に観測された、TRMMの可視赤外観測装置(VIRS)のch 1(0.63µm)に赤、ch 2(1.6µm)に 緑、ch 4(10.8µm)に青を振り分け合成したチベット付近の画像です。ヒマラヤの山岳地帯の積雪域が赤系、水雲の領域は 黄色系、ヒマラヤの谷や高原地帯が青系統の色に良好に判別できています。水雲はch 1及びch 2の反射がどちらも強く温度 は低いために黄色になります。積雪域はch 1の反射率は大きいのですが、ch 2の反射率が小さく温度が低いために赤系統の 色になります。他の陸域は温度が高く反射率が小さいために、青みがかって見えます。図2は図1と同じ領域のch 1とch 2の比 (ch 1/ch 2)の白黒画像です。この比の値が大きいほど白くなります。水雲の領域とヒマラヤの積雪領域で白さに大きな違い が見られます。ch 1では氷と水で吸収の違いは小さいのですが、ch 2では氷による吸収のほうが大きいために反射が小さくなっ ています。このため比をとると、水雲と積雪の判別が良好にできます。図3は同じ日にブラジルの南で観測された、ch 4とch 5 (12µm)の輝度温度差の小さい領域に赤、大きい領域に緑、ch 4に青を振り分け合成した画像です。光学的に厚い客域は 赤系統に、光学的に薄い氷雲は緑系統に、温度の高い陸地と海は青系統の色に見えます。積雲型の雲は光学的に厚いの でch 4とch 5の輝度温度差が小さく赤系統に、巻雲型の雲は輝度温度差が大きいので緑系統の色に見えます。或4 は図3と同じ領域のch 4とch 5の輝度温度差が小さいので赤系統に、水蒸気量の多い領域は青系統の色に見えます。図4 は図3と同じ領域のch 4とch 5の輝度温度差が小さいので赤系統に、水蒸気量の多い領域は青系統の色に見えます。 雲であるか、晴天域かをch 4の温度やch 1の反射率で判別したのち、輝度温度差の情報を用いると雲の種類、水蒸気量や 水蒸気の吸収による減衰を補正した海面温度の情報を得ることができます。

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Figure 1 is a color composite RGB image (channel 1 [0.63µm], 2 [1.6µm] and 4 [10.8µm] for red, green and blue, respectively) over the Tibetan plateau observed by the TRMM Visible Infrared Scanner (VIRS) on Dec.21,1997. The snow-covered Himalayan mountains are reddish, water cloud areas are vellowish, and Himalayan valleys and plain areas are bluish. The larger reflectivity for both ch 1 and ch 2 makes the water cloud yellowish, the larger reflectivity for ch 1 and smaller reflectivity for ch 2 makes snow areas reddish. The valleys and plain areas are bluish due to their relatively warmer temperature. Figure 2 is a gray-scale image for the ratio of ch 1 and ch 2. The observed region and time are the same as in Fig. 1. The larger region is white. Water reflectances for the ch 1 and ch 2 wavelengths are similarly, but ice reflectance is smaller for the wavelength of ch 2 than that of ch 1. Therefore, the ratio of ch 1 / ch 2 is larger for snow covered areas than water cloud areas. Figure 3 is a color composite RGB image (red for smaller brightness temperature difference (BTD) between ch 4 and ch 5 [12µm], green for the larger BTD, and blue for the brightness temperature (TBB) at ch 4) over the south of the Brazil observed in the same day. Optically thicker cumulus clouds are reddish due to the smaller BTD, optically thin ice clouds are green and ocean and land are blue because of the warmer TBB. Regions with less water vapor are reddish due to the smaller BTD, and moist regions are bluish over the ocean because of the larger BTD and warmer TBB. Figure 4 is the gray- scale image of BTD. The observed region and time are the same as in Fig. 3. The larger BTD is shown in white. Optically thin cirrus clouds exhibit larger BTD, while optically thick cumulus clouds indicate smaller BTD. Over cloud free ocean areas, the BTD increases with the increase of water vapor. The BTD is useful for classifying clouds and deriving water vapor amount. The Sea Surface Temperature, which is corrected for water vapor absorption, can also be estimated using ch 4 and ch 5.